# A Highly Configurable Vortex Initialization Method for Tropical Cyclones

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## Motivation

•Data assimilation is the future. But, currently:

- •Observations (that are assimilated) are sparse.
- •Computationally expensive.
- •Complicated with a long learning curve.

•We wish to provide an alternative to existing bogus methods.

•GFDL – axisymmetric spin-up. Asymmetries provided by asymmetric component of previous forecast at new initialization time. Too complex.

•WRF – Idealized Rankine. Axisymmetric bogus data. Asymmetry provided by smooth environmental field. Too simple.

•New bogussing technique is:

•Highly configurable to match any vortex shape.

•Specify full three dimensional wind field to minimize adjustment period.

# Algorithm

#### Vortex Removal

- Vortex Addition
  - Radial Structure:
    - Modified Rankine Vortex
    - Willoughby Vortex
  - Vertical Structure:
    - Boundary Layer
    - Free Atmosphere

# Algorithm – Vortex Removal Largely follows Kurihara et al. (1995)



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#### Algorithm – Radial Structure

Modified Rankine Vortex

$$V(r) = V_{\max} \left(\frac{r}{R_{\max}}\right) \qquad r < R_{\max}$$
$$V(r) = V_{\max} \left(\frac{R_{\max}}{r}\right)^{\alpha} \qquad r > R_{\max}$$

Willoughby Vortex (Willoughby et al. 2006)



FIG. 2. A dual-exponential profile used to approximate the observed wind in Hurricane Diana on 11 Sep 1984. Here and subsequent shading indicates observed winds, and the darker curves indicate the fitted profiles.

$$V(r) = V_1 = V_{\max} \left(\frac{r}{R_{\max}}\right)^n$$

$$V(r) = V_1 \left(-w\right) + V_o w.$$

$$V(r) = V_o = V_{\max} \left[\left(-A\right) \exp\left(-\frac{r-R_{\max}}{X_1}\right) + A \exp\left(-\frac{r-R_{\max}}{X_2}\right)\right]$$

$$r \le R_1$$
$$R_1 \le r \le R_2$$
$$R_2 \le r$$

**Configurable parameters:** 

Inner radial structure -  $V_{max}$  and  $R_{max}$ Outer radial structure -  $\alpha$  and  $X_2$ 

## Algorithm – Radial Structure



# Algorithm

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#### Algorithm – Vertical Structure: Boundary Layer

Boundary Layer flow follows Foster (2009):

Steady state, height dependent, axisymmetric flow under a specified wind field.

$$\begin{split} &\frac{\partial U}{\partial r} + \frac{U}{r} + \frac{\partial W}{\partial z} = 0. \\ &U\frac{\partial U}{\partial r} - \frac{V^2}{r} + W\frac{\partial U}{\partial z} - fV = \frac{-1}{\rho_o}\frac{\partial P}{\partial r} + \frac{\partial}{\partial z}(K\frac{\partial U}{\partial z}). \\ &U\frac{\partial V}{\partial r} - \frac{UV}{r} + W\frac{\partial V}{\partial z} + fU = \frac{\partial}{\partial z}(K\frac{\partial V}{\partial z}). \\ &K\frac{\partial \langle V, V \rangle}{\partial z} = \frac{\tau}{\rho_o} = C_D |\vec{V}|(U, V). \end{split}$$

Configurable parameters: Boundary layer height and constant eddy diffusivity, K

#### Algorithm – Vertical Structure: Boundary Layer



U

V

















# Algorithm

Vortex Removal

#### Vortex Addition

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#### – Vertical Structure:

- Boundary Layer
- Free Atmosphere

## Algorithm – Vertical Structure: Free Atmosphere Gaussian Decay







**Configurable parameters:** Altitude of maximum tangential wind,  $Z_{max}$ Decay parameters,  $L_{up}$ ,  $L_{down}$ ,  $\alpha_1$ , and  $\alpha_2$ 

## Algorithm – Vertical Structure: Free Atmosphere Emanuel Theory (1986)

 $R_{max} = R_{max}(z)$  using conservation of saturated moist static energy above the boundary layer.

• $V_{max}$ = $V_{max}(z)$  by solving for V at  $R_{max}$  noting conservation of angular momentum.

V(r, z) calculated from V(r) profile at each altitude above the boundary layer.



Configurable parameters: Boundary layer height

Outflow temperature (controls height of the vortex)

#### Algorithm – Vertical Structure: Matching

•Absolute angular momentum M(r, z) calculated from V(r,z).

• $\Psi(r)$  is calculated at the boundary layer top through the inward integration of vertical motion. Thus a functional relationship between  $\Psi$  an M (or  $\Psi(M)$ ) is determined.

• $\Psi(r, z) = \Psi(M)$ . Maintain constant  $\Psi$  along angular momentum surfaces as angular momentum is conserved above the boundary layer.

•U(r, z) and W(r, z) determined from  $\Psi(r, z)$ .



### **Testing - Real**

- WRF-ARW 3.1.1.
- 3 Grids (27/9/3 km).
- 40 vertical levels stretched in height.
- YSU boundary layer parameterization.
  - Modified drag formulation (Donelan et al 2004; Davis et al. 2008).
- WRF 6-species microphysics (single-moment).
- RRTM longwave and Goddard shortwave parameterizations
- Grell-Devenyi ensemble cumulus package on outermost grid.

## Modified Rankine (No SC) vs. Willoughby (SC)















U

V





U

V



#### Real: ModRank (V) vs. Willoughby (UVW) – Hour 4



U

V





U

V





U

V





U

V





U

V



## Testing - Idealized

- WRF-ARW 3.1.1.
- Constant SST =  $28.5 \, {}^{\circ}$ C.
- No-SAL Jordan sounding.
- No environmental flow.
- 3 Grids (27/9/3 km).
- 40 vertical levels stretched in height.
- No radiation/convection parameterization.
- YSU boundary layer parameterization.
  - Modified drag formulation (Donelan et al 2004; Davis et al. 2008).
- WRF 6-species microphysics (single-moment).

#### Ideal: Modified Rankine (No SC) vs. Willoughby (SC)













U

V





U

V





U

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U

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#### Willoughby (SC) vs Willoughby (SC-No Mass Pert.)





U

V





U

V





U

V





U

V





U

V



### Conclusions

- A highly configurable vortex initialization methodology has been constructed that allows precision manipulation of the initial vortex structure.
- The configuration options range from the highly simplistic to the highly complex in which a continuous boundary layer/free atmosphere vortex flow with a mass conserving secondary circulation may be implemented.
- Several test cases show that initial spin-down of the vortex, from a structural perspective, is reduced when the full three dimensional wind field is accounted for.